A large, semi-transparent image of fossilized bird wings is centered in the background of the slide. The wings are dark and show the structure of the feathers and bones.

# Climate reconstruction from fossil data using a Bayesian approach

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## Outline

- Introduction
  - Data
  - Model
  - Results
  - Conclusions
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- A vertical strip on the right side of the slide shows a close-up of a fossilized bird wing, similar to the one in the top slide, but smaller and partially cut off.

## Starting point

- Forecasting climate change requires information on the past climate behavior
- Meteorological records are too short to cover the full climate variation
  - > History has to be inferred from indirect “proxy” indicators
- Biological indicators preserved in the sedimentary deposits are a major source of proxy-climate data

## Reconstruction

- Early attempts for climate reconstruction utilized qualitative and descriptive analysis on single indicator species data
- Nowadays numerical techniques and approaches are available allowing quantitative reconstruction
- In general, the methods involve two steps
  - Regression: mapping responses of biological indicators to the contemporary environment
  - Calibration: predicting the environmental variable from the fossil data

## Approaches

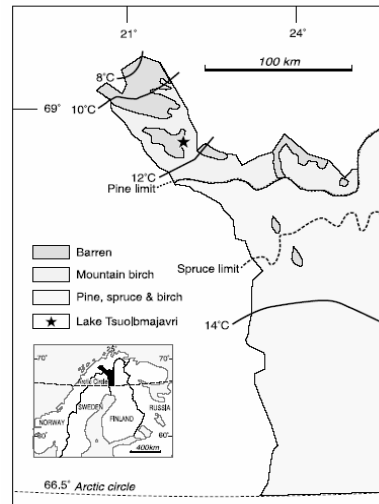
- Ways to perform the quantitative reconstruction can be classified to two categories:
  - Classical: estimate the response function from the climate state to the biological indicator, and reconstruct the climate state by using the inverse of the function
  - Inverse: estimating directly the inverse function from the biological indicators to the climate and applying the function in the reconstruction
- Reconstruction contains many uncertainties to be accounted for
  - Nature in general is not deterministic
  - Measured biological indicator data is “noisy”
  - Applied model only approximates the real world

## Objective

- Applying Bayesian multinomial Gaussian response model to fossil chironomid assemblages in a tree-line lake in Finnish Lapland to reconstruct Holocene palaeotemperatures
  - All unknown quantities – parameters as well as the data – are treated as random variables
  - The gained knowledge and reconstructions are explicitly conditioned on the statistical model, prior knowledge and data in hand

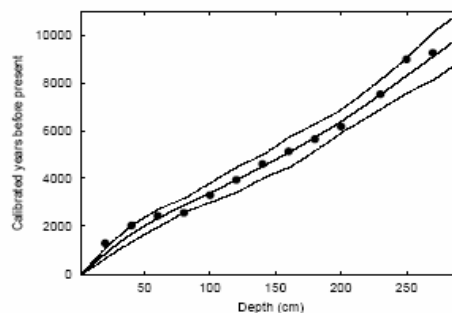
## Study site

- Lake Tsuolbmajavri is located on the north-western Finland, 526m above sea level at the boundary between boreal forest and tundra
  - It is small ( $A=13.9$  ha), shallow ( $Z_{\max}=5.35$  m), clear and oligotrophic
  - Measured mid-summer water temperature is  $10.9^{\circ}\text{C}$  and site-specific altitude-corrected mean July air temperature is  $11.0^{\circ}\text{C}$ .
  - The mean annual air temperature is about  $2^{\circ}\text{C}$ ; annual precipitation is about 350mm (50% falls between June and September)



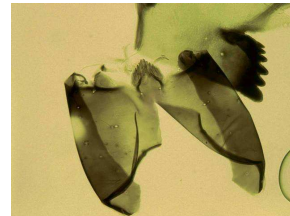
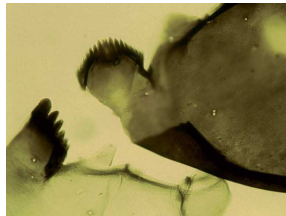
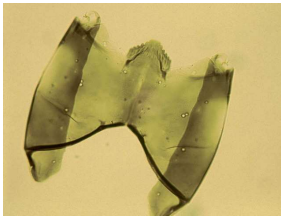
## Fossil data

- A 291-cm-long sediment core was sampled in 1997 at the central part of the lake from the frozen lake surface
- For chronology 14 radiocarbon ( $^{14}\text{C}$ ) datings were determined from the core and age-depth models were developed by non-parametric weighted regression
- The dates suggest a fairly linear sediment accumulation rate



## Chironomid analysis

- The sediment core was sampled for subfossil chironomids at intervals of 2 cm, which is equivalent to a temporal resolution of 50–70 years
- Subsequently, a few samples were taken at 1 cm interval from sequences with rapid changes in chironomid composition
- Chironomid head capsules were counted from the samples using complex, manual procedure (including chemical processing, sieving, microscoping etc.)
  - 50 chironomid taxa were recorded from the total of 148 samples

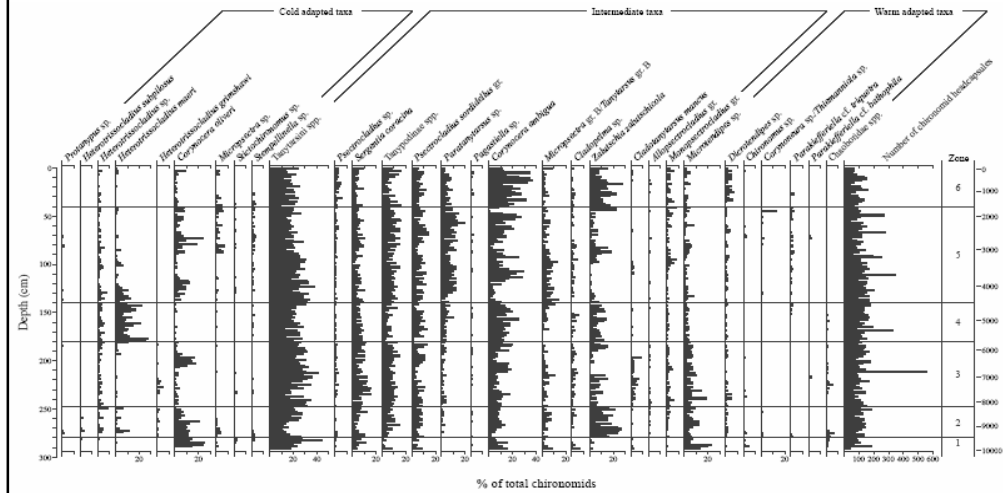


## Contemporary data

- The modern training data used for calibration models include surface-sediment chironomid head capsules and air temperature data from 63 lakes in north-western Finland.
  - The study lakes are generally small, clear and oligotrophic
- Chironomid data are expressed as the head capsule count
  - Rare taxa (less than 2% relative abundance in more than 4 lakes) were excluded
  - The remaining data included 52 chironomid taxa, 45 of which are common with the fossil data
- Mean July temperatures were estimated for each lake using data of nearby climate stations from years 1961-90
  - One measurement was removed as an outlier
  - The remaining data included 62 measurements ranging from 7.9°C to 13.8°C



## Chironomid stratigraphy



## Problem statement

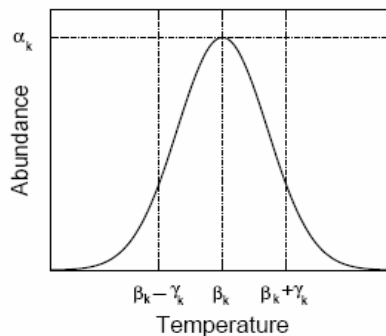
- Let  $y_{ik}$  be the count of taxon  $k$  at site  $i$  and  $y_i = (y_{i1}, \dots, y_{im})^T$  be the actual counts at site  $i$   
 ->  $Y = [y_1, \dots, y_n]$  is the  $m \times n$  matrix of abundances ( $m=52$ ,  $n=63$ )
- Furthermore let this be denoted as  $\tilde{y}_{ik} = y_{ik}/y_{i+}$ , where  $y_{i+} = \sum_h y_{ih}$  (relative abundance)
- Let  $x_i$  be the temperature at site  $i$   
 ->  $X = (x_1, \dots, x_n)^T$  is the vector of temperatures at sites
- Let  $y_0 = (y_{01}, \dots, y_{0m})^T$  be a fossil assemblage and  $x_0$  the air temperature at the time the fossil assemblage was formed
- Given  $Y$ ,  $X$  and  $y_0$ , the task is to reconstruct the temperature  $x_0$ .
  - In other words the idea is to estimate the posterior distribution  $f(x_0|X, Y, y_0)$

## Bayesian approach

- Let  $\theta$  be the vector of parameters of the statistical model  
->  $f(x_0|X, Y, y_0) = \int f(x_0, \theta|X, Y, y_0) d\theta$   
 $\propto \int f(x_0, \theta) f(Y, y_0|X, x_0, \theta) d\theta$
- In other words, the model specifies the response functions that consist of two parts
  - Likelihood term  $f(Y, y_0|X, x_0, \theta)$
  - Prior probability distribution  $f(x_0, \theta)$

## Bummer model

- Bummer uses a multinomial response model  
->  $E(\tilde{y}_{ik}|x_i, \theta_k) = \lambda_{ik} / \sum_j \lambda_{ij}$ , where  $\lambda_{ik}$  is the response of taxon  $k$  to temperature at site  $i$
- It is assumed that the response curve has a unimodal Gaussian shape  
->  $\lambda_{ik} = \alpha_k \exp\{-[(\beta_k - x_i)/\gamma_k]^2\}$
- Hence  $\theta_k = (\alpha_k, \beta_k, \gamma_k)$  represents hyperparameters that control scaling factor, optimum and tolerance of taxon  $k$  to the temperature



## Likelihood term

- It is assumed that sites are mutually independent given the temperatures and taxon-specific model parameters  
->  $f(Y, y_0 | X, x_0, \theta) = \prod_i f(y_i | X, x_0, \theta) = \prod_i f(y_i | x_i, \theta)$
- It is further assumed that the taxon abundances  $y_i$  are multinomially distributed  
->  $(y_{i1}, y_{i2}, \dots, y_{im} | y_{i+}, p_i) \sim \text{Mult}(y_{i+}, p_i)$ , where  $p_i = (p_{i1}, \dots, p_{im})$  and  $p_{ik}$  is the probability of an individual from site  $i$  to be of taxon  $k$
- Random variables  $p_i$  are assumed to be Dirichlet distributed  
->  $(p_{i1}, \dots, p_{im} | x_i, \theta) \sim \text{Dirichlet}(\lambda_{i1}, \dots, \lambda_{im})$ , given the Gaussian responses  $\lambda_{ij}$

## Prior distributions

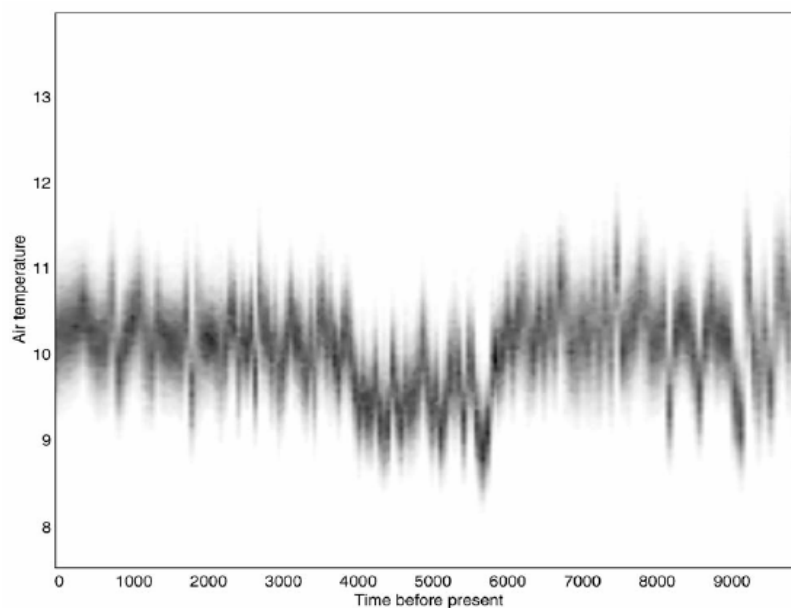
- It is assumed that each temperature  $x_0$  to be reconstructed as well as all species-specific parameters  $\alpha_k$ ,  $\beta_k$  and  $\gamma_k$  are mutually independent in the prior  
-> The joint prior distribution  $f(x_0, \theta)$  can be determined by individual marginal priors
- The following parameter *a priori* distributions were chosen
  - $x_0$  is normal distributed with mean 11 and variance 1
  - $\alpha_k$  is uniformly distributed in [0.1, 50]
  - $\beta_k$  is normal distributed with mean 11 and variance 3
  - $\gamma_k$  is gamma distributed with mean 3 and variance 1



## Modeling

- The integration of the model was implemented using stochastic Markov chain Monte Carlo (MCMC) methods
- Statistical accuracy of the reconstructed temperature was assessed using the contemporary data
  - 0.8°C root mean square error of the prediction calculated by leave-one-out cross-validation
  - Maximum bias along the temperature gradient 0.98°C
- Bummer has high predictive and low inherent systematic and maximum bias
  - > Model should be sufficiently precise for tracking temperature changes, given the assumed sensitivity of the proxy source

## Inferred temperature



## Reliability of the reconstruction

- There are many uncertainties and potential problems to be taken into account in the interpretation
  - Many of finer scale variations are smaller than the average accuracy of the model
  - Chronological control and chironomid analysis involve uncertainties
  - Stochastic nature of the environment
- Large-scale patterns were assessed by comparing with independent proxy records from the same region
  - > Results seem accurate enough, apart from the late-Holocene (due to shallowing of the lake by progressive sedimentation)
- The uncertainty in the fine-scale changes was minimized by
  - Using taxa whose ecological indicator value is well-established
  - Ensuring that the taxa were well represented in fossil and contemporary data
  - Clarifying the mechanisms and causations behind the changes

## Interpretation of the reconstruction

- Climate during the early Holocene was particularly unstable
  - However, concerning the very first few hundreds of years of the Holocene the prediction accuracy (credibility) is quite low
- Fairly low but steadily rising summer temperatures during the early Holocene.
  - There are, however, three successive cooling events with amplitudes of 1–1.5°C at around 9200, 8600 and 8300 cal yr BP
- The interval between ca 8000 and 5800 cal yr BP appears to have been warm and stable
- Relatively long-lasting temperature minimum between ca 5800 and 4000 cal yr BP (ca 1°C drop in the inferred mean)
  - Within the cold interval, particularly cool summers appear to have prevailed at about 5800, 5000 and 4200 cal yr BP
- The inferred late-Holocene climate is surprisingly featureless, fluctuating between 10–10.5°C during the past 4000 years.
  - A brief cold oscillation can, however, be distinguished with some confidence at about 1800 cal BP

## Summary

- A Bayesian model was implemented for reconstructing long-term temperature changes from subfossil chironomid remains
  - Representing the temperatures as posterior distributions involves explicit information on the model statistical uncertainty
  - Embedded ecological knowledge results in more plausible model
  - Classical approach enables improved extrapolation capability
- Obtained summer temperature reconstruction is in general consistent with other independent proxy sources from the region
  - Unstable early Holocene, cooling events at 9200, 8600 and 8300 cal yr BP, thermal maximum ca 8000-5800 cal yr BP, and distinct cooling around 5800 cal yr BP
  - Late-Holocene temperature is apparently not reconstructed accurately due to the shallowing of the lake by progressive sedimentation

## References

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- Toivonen, H.T.T., Mannila, H., Korhola, A., Olander, H., 2001. *Applying Bayesian statistics to organism-based environmental reconstruction*. *Ecological Applications* 11, 618–630.
- Korhola, A., Vasko, K., Toivonen, H.T.T., Olander, H., 2001. *Holocene temperature changes in northern Fennoscandia reconstructed from chironomids using Bayesian modelling*. *Quaternary Science Reviews* 21, 1841–1860